

SPACE Geodesy and NASA data in Altimeter Satellite POD¹

POD: Precise Orbit Determination



Altimeter satellites

Frank G. Lemoine Planetary Geodynamics Laboratory, Code 698 NASA Goddard Space Flight Center June 29, 2016



Precise Orbit Determination

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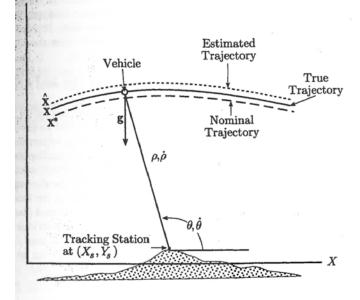


Figure 1.2.1: Uniform gravity field trajectory.

"Statistical Orbit Determination" BD Tapley, BE. Schutz, GH Born, Elsevier Academic Press, 2004. Orbit Determination: The problem of determining the "best estimate of a satellite ephemeris (orbit) using observations that are influenced by random and systematic errors, using mathematical models of spacecraft motion that are not exact.

<u>What is meant by "precise"?</u> In the past 25 years it has come to mean with a radial RMS accuracy of a few cm or

<u>Why?</u> Because the analysis of our science observations requires a "reference" that is accurate to this level; Sea level; Natural Hazards (Crustal Deformation); Glacier or Ice Sheet height changes.



Motivation for POD - I



GEOS-3

(Geodynamics and Earth Ocean Satellite)

Launched: Apr. 9, 1975 Operated through July 1979.

B. Douglas et al., JGR, 1983, http://dx.doi.org/10.1029/JC088iC14p09595

GEOS-3 COLLINEAR ALTIMETER DATA

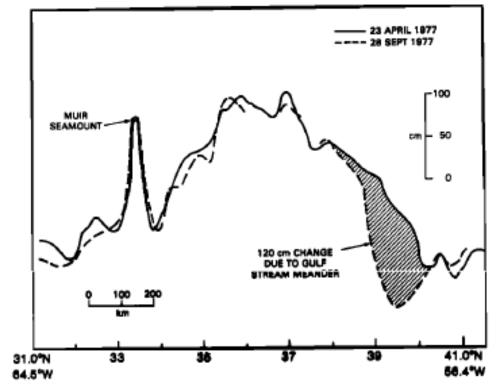


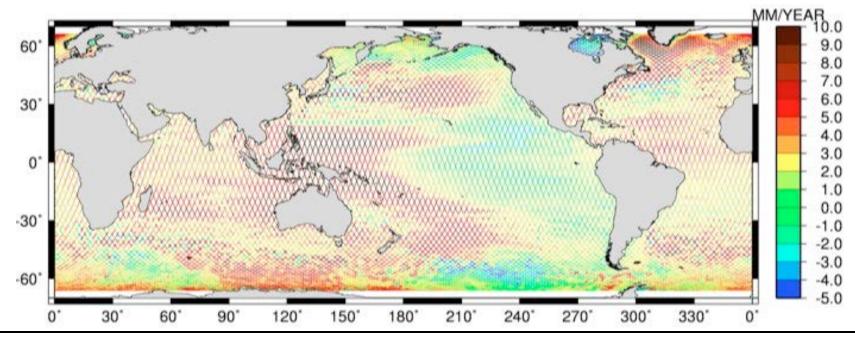
Fig. 3. A close collinear pair of altimeter profiles crossing the Muir seamount north of Bermuda. Note the identical geoid undulation in the profiles at the position of the seamount. The shaded area is due to a meander of the Gulf Stream.

An altimeter measures ocean height, which provides information on the geoid, and on the changes in the ocean circulation.

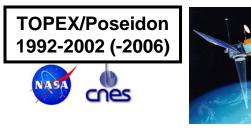


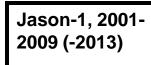
Motivation for POD - II

Measurement of Regional and Global Mean Sea Level Change



Regional mean sea level variations from TOPEX, Jason-1, and Jason-2 with respect to 1993-2002 mean; http://podaac.jpl.nasa.gov/Integrated_Multi-Mission_Ocean_AltimeterData





Jason-2, 2008-Jason-3, 2016-





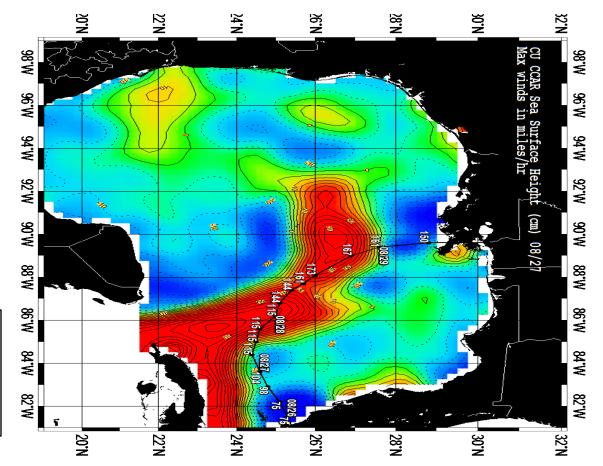


Motivation for POD - III

Hurricane Intensification from Passage over Warm Core Eddies

Sea Surface Height variations show the location of warm water eddies – which appear higher in absolute height. Their latent height can contribute to hurricane intensification.

Mapping of Gulf of Mexico Sea Surface Height Variations by Dr. Robert R. Leben, University of Colorado, Boulder.



http://oceanmotion.org/html/impact/natural-hazards.htm http://www.nasa.gov/centers/jpl/news/ostm-20080701.html



Current Ocean-radar mapping altimeter satellites (June 2016)



SARAL, 2013-



Haiyang (HY)-2A, 2011 (CNSA)

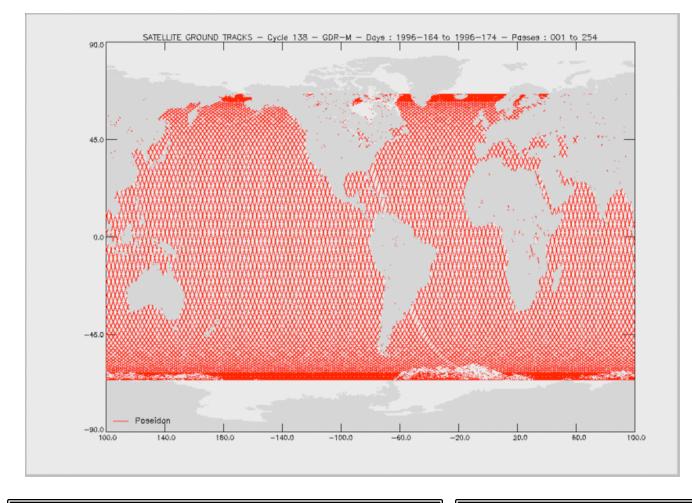




Sentinel-3A 2016 From the launch of the first spaceborne altimeters, Precision Orbit Determination (POD) has been driven by the science goals of the geodetic altimeter missions...



Example Ground Track Coverage for TOPEX (& Jason-1, Jason-2, Jason-3)





TOPEX/Poseidon 1992-2006



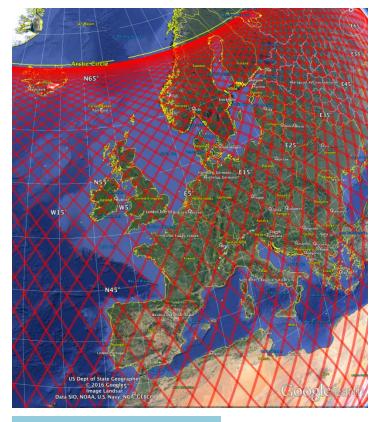
Jason-3 2016 –

Image from AVISO (Toulouse, France)

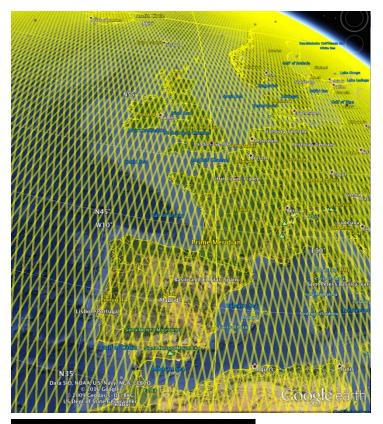
Altitude 1336 km. Inclin. = 66.039°; Ground track repeat: 9.9156 days. Cross-track separation (equator): 315 km



Example: Ground Track Coverage for TOPEX vs. ERS/Envisat



TOPEX/Jason-1,2,3

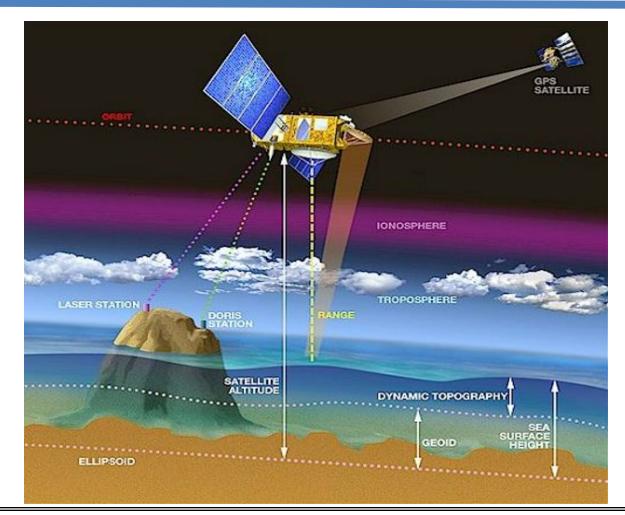


ERS & Envisat & SARAL

Altitude ~785 km. Inclin. 98.543°; (sun-synchronous) Ground track repeat: 35 days. Cross-track separation (equator): 80 km



POD - Schematic

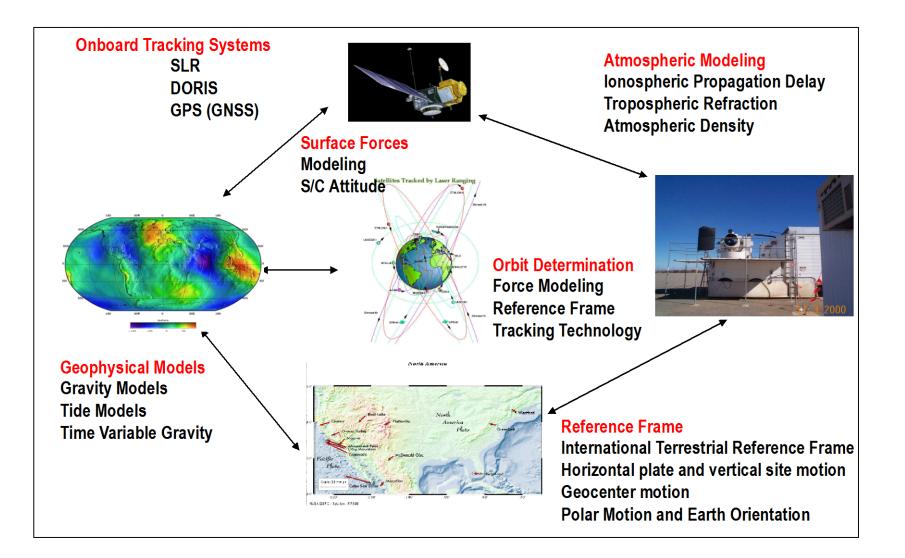


This example is for TOPEX, but the same principle applies for all altimeter satellites.

In order to determine the height of the sea surface, we must know the satellite position (meaning its orbital ephemerides) to a precision commensurate to or better than the accuracy of the altimeter

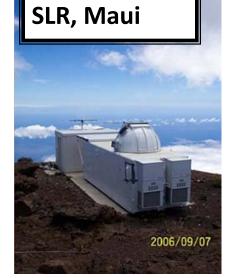


Orbit Determination Schematic



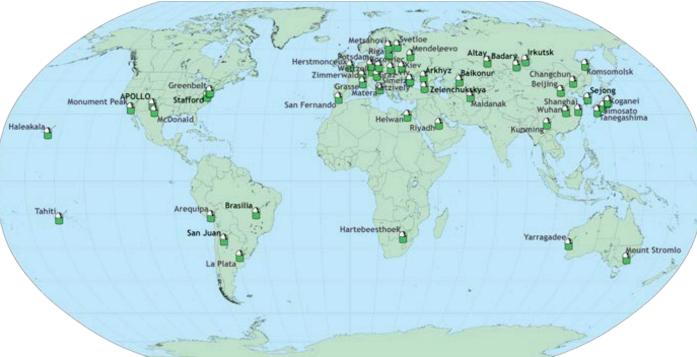


Satellite Laser Ranging (SLR)





SLR, Yarragadee, Australia



- Wavelength: 532 nm.
- Best stations have mm precision.
- Preponderance of stations in N. Hemisphere.
- First satellite ranging 1964, NASA GSFC.

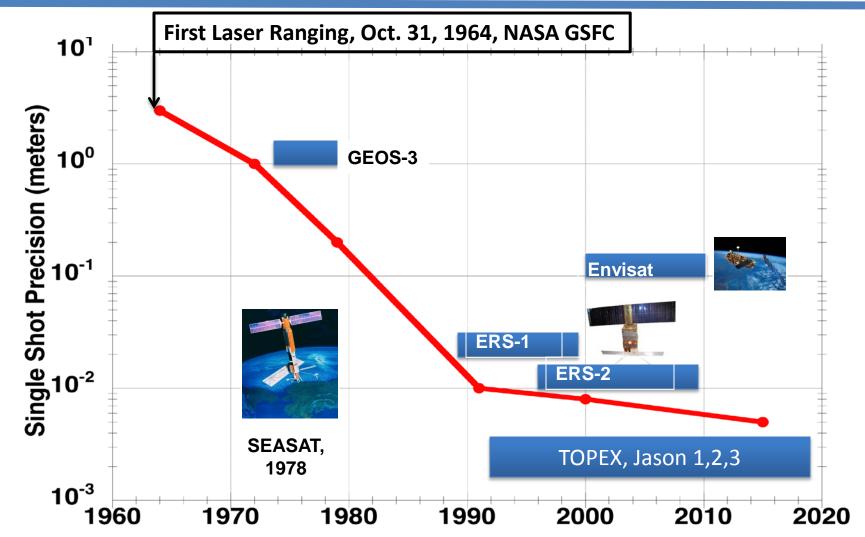
http://ilrs.gsfc.nasa.gov

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International Laser Ranging Service



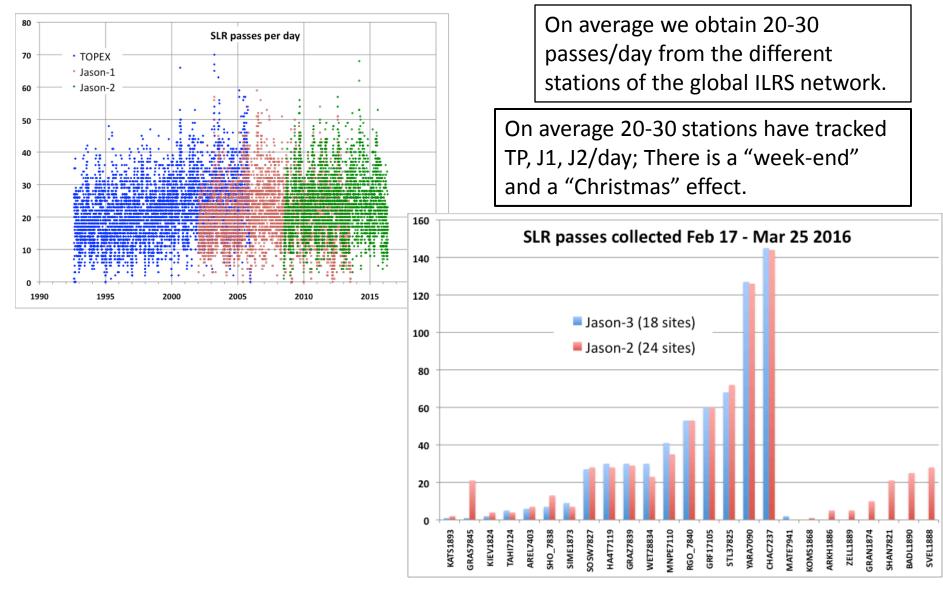
Representative SLR precision vs. time



Adapted from J. Degnan. "Impact of SLR Technology Innovations on Modern Science", 18th ILRS Workshop, Fujiyoshida, Japan, Nov. 11, 2013. http://cddis.gsfc.nasa.gov/lw18/docs/presentations/Session0/13-0001-Degnan 2.pdf



SLR Tracking of TOPEX, Jason-1 & Jason-2





DORIS (Doppler)

The satellite is upright the beacon, it's the DORIS: Doppler Orbitography and TCA point (Time of Closest Approach). The frequency of the received signal is equal Radiopositioning Integrated by Ranging. to the frequency of the transmitted signal. • CNES (French Space Agency) & IGN (Institut Géographique National) The satellite is approaching the The satellite is moving away the beacon: beacon: • 55-60 stations, around the world. The frequency of the received signal The frequency of the received is greater than the frequency of the signal is **lower** than the frequency Dual-frequency beacons, transmitted signal. of the transmitted signal. ~401.25 MHz + 2.036 GHz. DORIS receiver: 1 channel (1990's); 2 channels (Envisat & Jason-1, SPOT-5); Doris Receiving-7 channels (Jason-2 and all later satellites) Antenna 1991 1995 1999 2003 2007 2011 2019 Transmitting SPOT2 beacon $\bigcirc \bigcirc \diamondsuit$ **Topex-Poseidon** ●◎≧ SPOT3 SPOT4 000 •••• Jason-1

Envisat 0Past missions 0.000 SPOT5 **Current missions** 0.0000 Jason-2/OSTM **Future missions** ●<</p>
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< Cryosat-2 Orbit determination \mathbf{O} HY-2A HY-2B Earth gravity field ● C Saral/AltiKa Earth rotation ●<</p>
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< Jason-3 Positioning ● COO Sentinel-3A Sentinel-3B COn-board orbit determination ● COO Jason-CS ⊕Time tagging

Images Credit:. Centre National d'Etudes Spatiales (CNES) Collecte Localisation Satellites (CLS)



DORIS Network (as of June 2016)



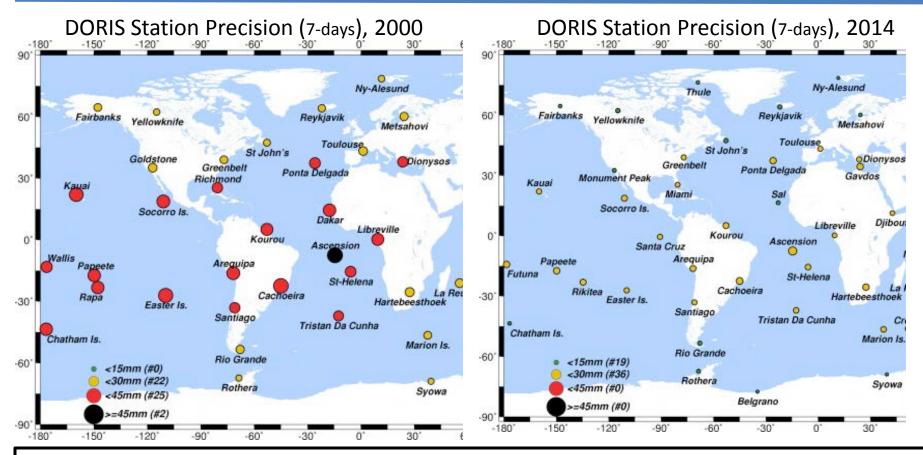


http://ids-doris.org/

International DORIS Service



DORIS Network Evolution



DORIS Network station positioning precision has improved in time, due to (1) improvements monumentation stability, (2) increasing number of satellites, (3) DORIS receiver improvements (more channels).

Reference: Guilhem Moreaux et al., Adv. Space Res., In press, 2016, doi: 10.1016/j.asr.2015.12.021

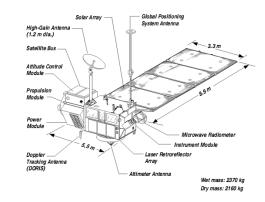
See also. Fagard, H., "Twenty years of evolution for the DORIS permanent network", J. Geodesy, 80 (8–11), (2006), pp. 429–456, doi: 10.1007/s00190-006-0084-2.



GPS Tracking System for TOPEX, Jason-1, Jason-2, Jason-3 – I



http://www.igs.org For latest information on IGS, see website and proceedings of IGS workshop, Sydney Australia, February 2016 TOPEX GPS: Demonstration Receiver



JASON 1-2-3: GPS Receiver

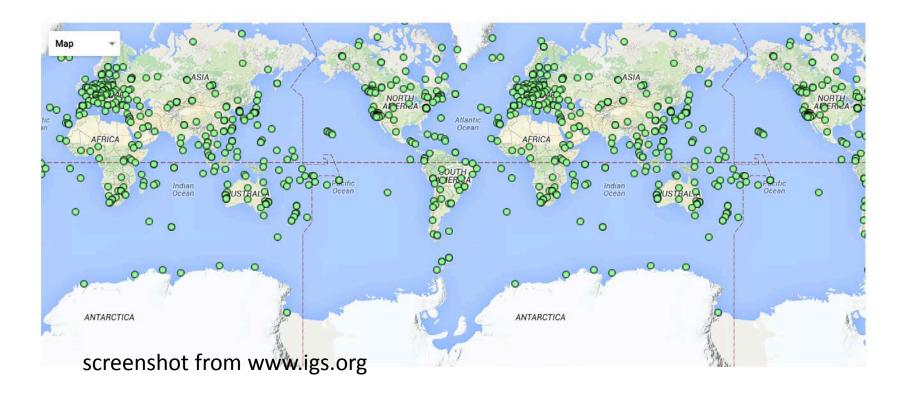


Examples: Ground Receivers









• The IGS includes hundreds of stations that stream data to the IGS data centers, including the NASA Crustal Dynamics Data Information System (CDDIS).

• Analysis centers around the world analyze the data over different latencies and deliver products (e.g. precise orbits) to users – the NASA CDDIS is one of the archives for these data products.





From W. Bertiger et al.,

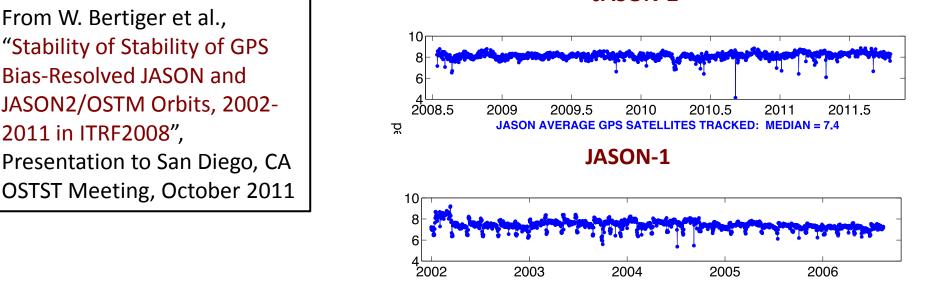
"Stability of Stability of GPS

JASON2/OSTM Orbits, 2002-

Bias-Resolved JASON and

2011 in ITRF2008",

GPS Data Quality Over Time (J1 & J2) **JASON-2**



The primary advantages of GPS as a (spaceborne) tracking system are the data density (# of observations per epoch) and geometrical strength of the observations.



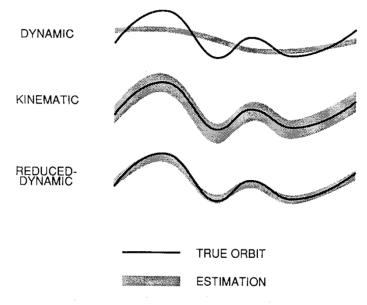


Fig. 1 Qualitative comparison of dynamic, kinematic, and reduceddynamic tracking performances.

Reduced-Dynamic Technique for Precise Orbit Determination of Low Earth Satellites

S. C. Wu,* T. P. Yunck,† and C. L. Thornton‡ Jet Propulsion Laboratory, California Institute of Technology, Pasadena, California 91109

J. Guidance & Control., 1991

Data density from GNSS, lately also from DORIS allows a reduced-dynamic orbit determination – which means empirical parameters can be adjusted frequently to reach the "true" orbit – and compensate for errors in force or measurement models.

Comparison of red.-dynamic & dynamic orbits allows insight into model errors – leads to improvement into dynamic models.



We need multiple tracking systems (a) to ensure and establish orbit accuracy;

This is especially important for the demanding application of measurement of the change in global mean sea level & to demonstrate orbit accuracy.

(b) to ensure redundancy; in the event one tracking system has "problems", or even fails.

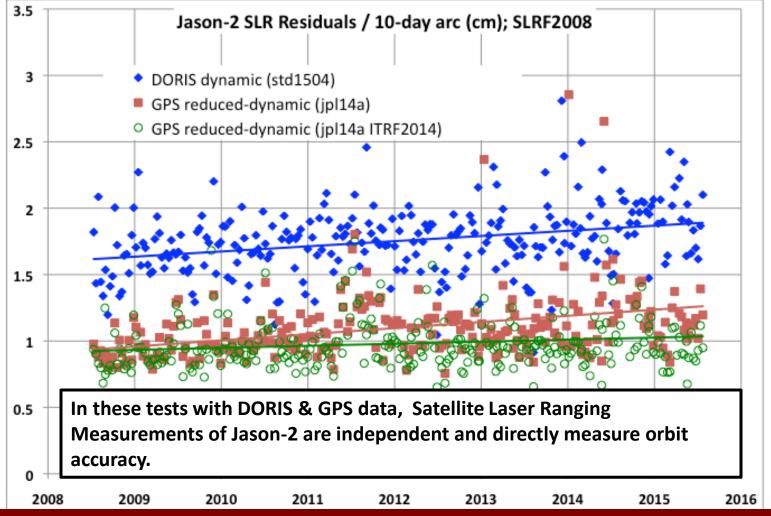
(I) GFO. Failure of GPS. SLR + altimeter crossovers only reliable tracking system.

(II) Jason-1. DORIS Oscillator not hardened before launch – perturbed by passage through S. Atlantic anomaly, Apply a "correction" model.



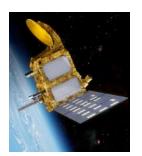
SLR – Validation of Jason-2 GPS & DORIS orbits

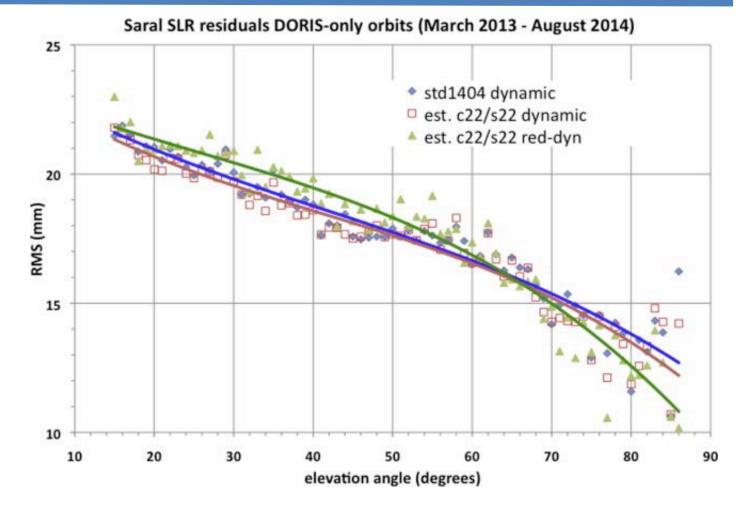




The fact that these orbits from different tracking systems agree at ~1cm radial RMS, is a reason why we can have such high confidence in the determination of Mean Sea Level change from satellite altimetry.

SLR – Evaluation of DORIS-only orbits (Saral)



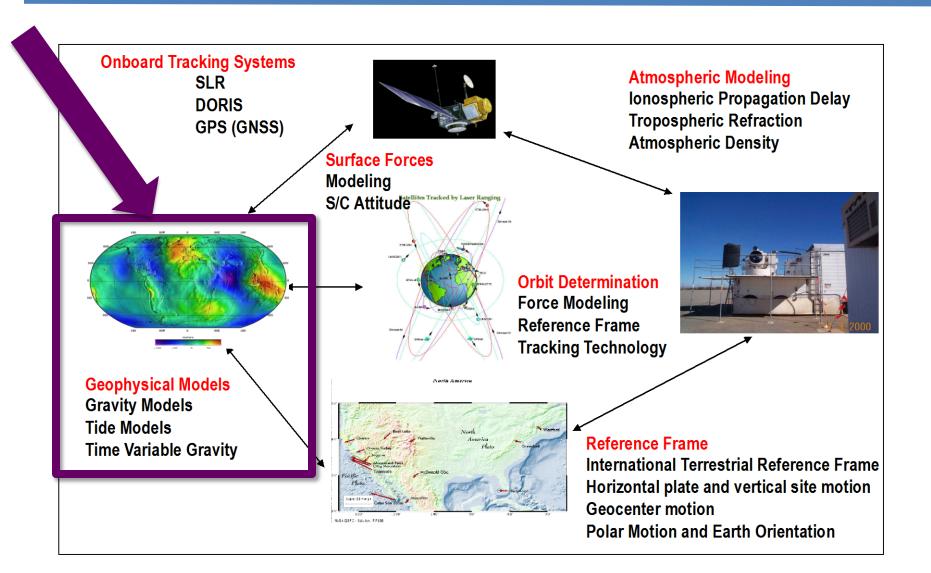


At high elevations SLR measures directly the radial orbit error; So in this example, we can say the DORIS-only orbits on SARAL have an orbital accuracy of 10-15 mm.

(Zelensky et al., 2016, "Towards the 1-cm SARAL orbit", Adv. Space Res, doi: 10.1016/j.asr.2015.12.011)



Orbit Determination Schematic

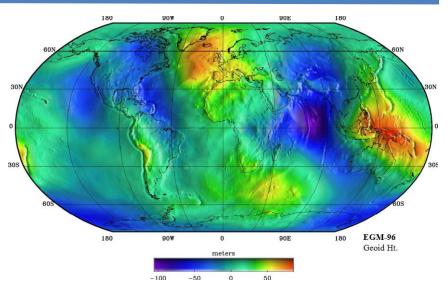




Representation of Gravity Field of the Earth (Spherical Harmonic Expansion) (Heiskanen & Moritz, 1967; Kaula, 1966)

$$U = \frac{GM}{r} + \frac{GM}{r} \sum_{\ell=2}^{\infty} \sum_{m=0}^{\ell} \left(\frac{R_e}{r}\right)^{\ell} \overline{P}_{\ell m}(\sin \theta)$$
$$\left(\overline{C}_{\ell m} \cos(m\varphi) + \overline{S}_{\ell m} \sin(m\varphi)\right), \qquad (1)$$

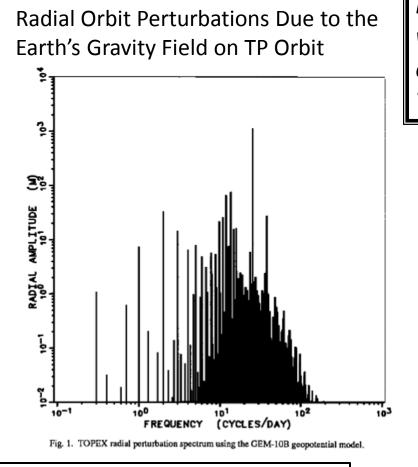
where G is the gravitational constant, M is the mass of the Moon, $\overline{P}_{\ell m}$ are the normalized associated Legendre polynomials of degree ℓ and order m, R_e is the reference radius (6378 km), and φ , θ , and r are the longitude, latitude, and radius at the evaluation point. $\overline{C}_{\ell m}$ and $\overline{S}_{\ell m}$ are the normalized Stokes coefficients, the main parameters of interest of our work.



EGM96 Geoid (Surface of constant equipotential)



Gravity Model Impact on Orbit Determination-II



Errors in Models of the Earth's Gravity Field were the largest source of orbit error for altimeter missions ... Until the launch of TOPEX/Poseidon

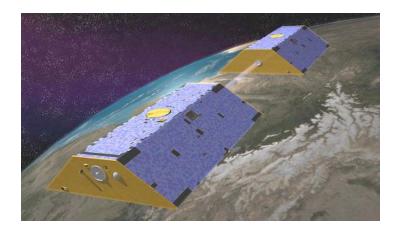
Model	L max x Mmax	SLR RMS of fit (cm)
GEML2, 1982	20x20	105.9
GEMT1, 1988	36x36	31.4
JGM-1S, 1991	70x70	7.7
JGM-2S, 1992	70x70	4.0
JGM-3, 1995	70x70	3.2
EGM96, 1996	70x70	2.8

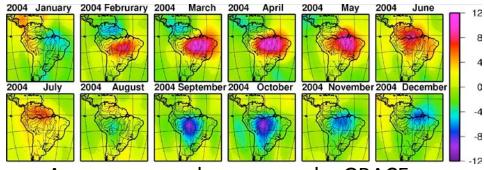
(BD Tapley & GW Rosborough, "Geographically Correlated. Orbit Error and Its Effect on Satellite Altimetry Missions" J. Geophys. Res., 1985, doi:10.1029/JC090iC06p11817)

The latest gravity models derived from GRACE & GOCE data eliminate static gravity error on the TP (J1, J2, J3) orbit and allow us to model in detail the temporal gravity variations



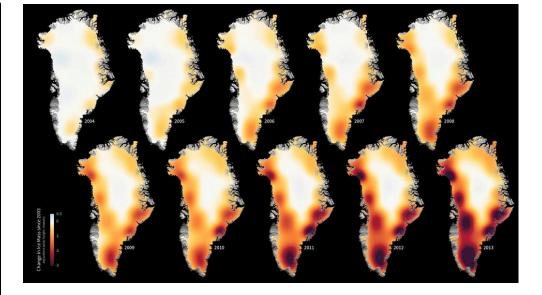
Gravity Model Impact on Orbit Determination-IV





Amazon mass change seen by GRACE http://grace.jpl.nasa.gov

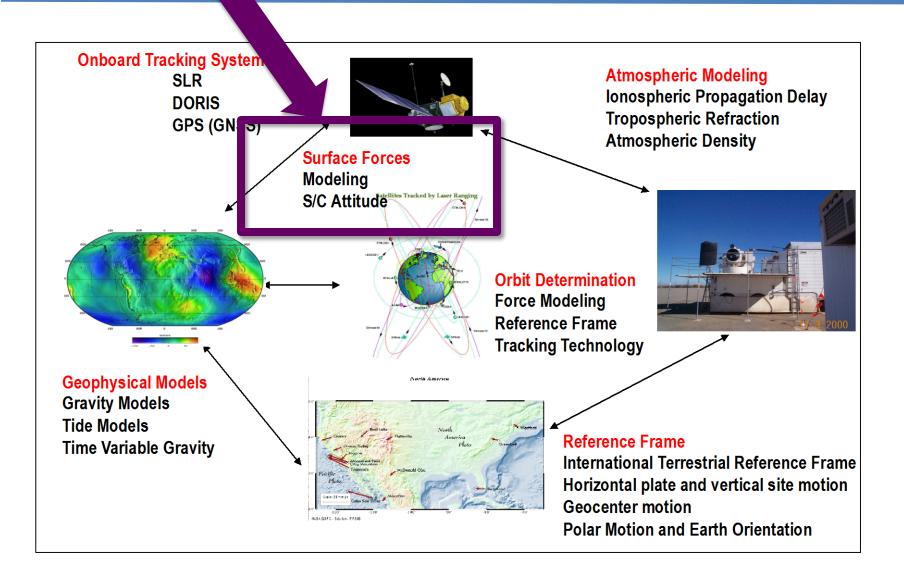
- US (NASA/JPL/UT CSR) + Germany (DLR),
- Launched 2002.
- Reliable 10-day to monthly solutions since January 2003 to present; Three official analysis centers (Univ. Texas, JPL, GFZ (Germany)); Also GRGS/CNES, NASA GSFC ...
- Most recent data has periodic gaps due to aging of s/c.



Mass loss in Greenland (2003-2013) https://svs.gsfc.nasa.gov/30478



Orbit Determination Schematic





Magnitude of Perturbing Forces on GEOSAT Follow On

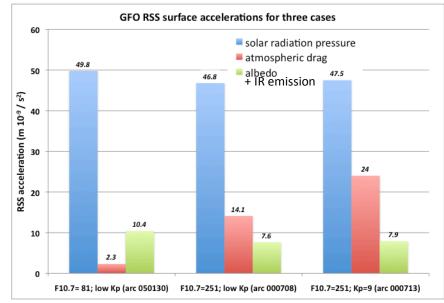
Perturbing Acceleration	m/s²	
GM	~7.7	
C _{2,0}	~9 x 10 ⁻³	
C _{20,20}	~1 x 10 ⁻⁷	
Ocean tides	~3 x 10 ⁻⁶	
Moon (third body)	~1 x 10 ⁻⁶	
Sun (third body)	~6 x 10 ⁻⁷	
Relativity – Schwarzschild	~1.4 x 10⁻ ⁸	
Solar Radiation Pr.**	~10 ⁻⁸	
Atmospheric Drag **	~10 ⁻⁷ to 10 ⁻¹⁰	
** Doponds on Aroa/Mass ratio 8		

** Depends on Area/Mass ratio & Position in solar cycle

Impact of Surface Forces depend on spacecraft shape & orientation, and for drag, on timing w.r.t solar cycle



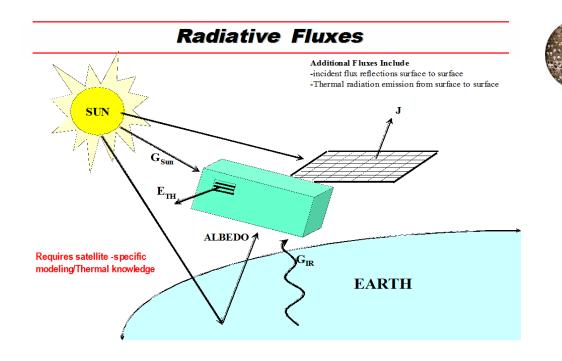
GFO-1, 1998-2008



GFO Surface Force Accelerations calculated by N. Zelensky (SGT @ NASA GSFC)

NASA

Radiation Pressure Modelling is the largest source of orbit error after gravity model error And remains a challenge

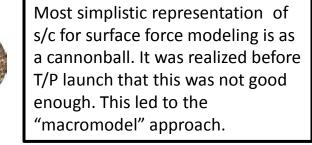


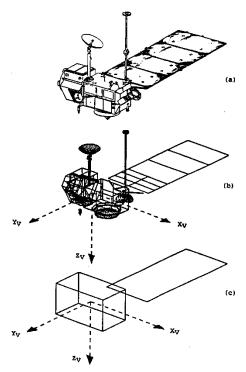
Micromodel:

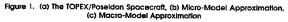
(Antreasian, 1992; Antreasian & Rosborough, 1992)

Box-Wing model

(Marshall & Luthcke. 1994)



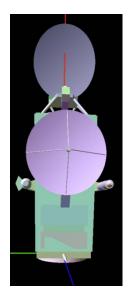






<u>One example:</u> University College London models for LEO spacecraft. (*Ziebart, 2004; Ziebart et al., 2005*)





Clean room photo (left) Mathematical model (right) A detailed s/c model with a thermal properties at ever node is illuminated from a simulated solar source over many orientations. Ray-tracing is used to "precalculate" the radiation-pressure accelerations as a function of spacecraft orientation.

These calculations are computationally intensive. We only initiate them if the mission requirements demand the highest orbit accuracy.

The extra effort can remove systematic signals in orbits & altimeter data.



Animation – Orbital Motion of Jason-2 Spacecraft

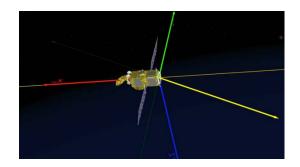
3D Animation of Jason-2 spacecraft in orbit provided by the International DORIS Service (IDS)

http://ids-doris.org/satellites.html

HY-2A, Cryosat-2, SPOT-5, Envisat also available.



International DORIS Service



<u>Conclusion</u>: We need detailed attitude information about orientation of spacecraft and solar arrays as a function of time in order to properly model the surface forces.

For the Jason s/c these are supplied as quaternions which the project archives on the NASA CDDIS;

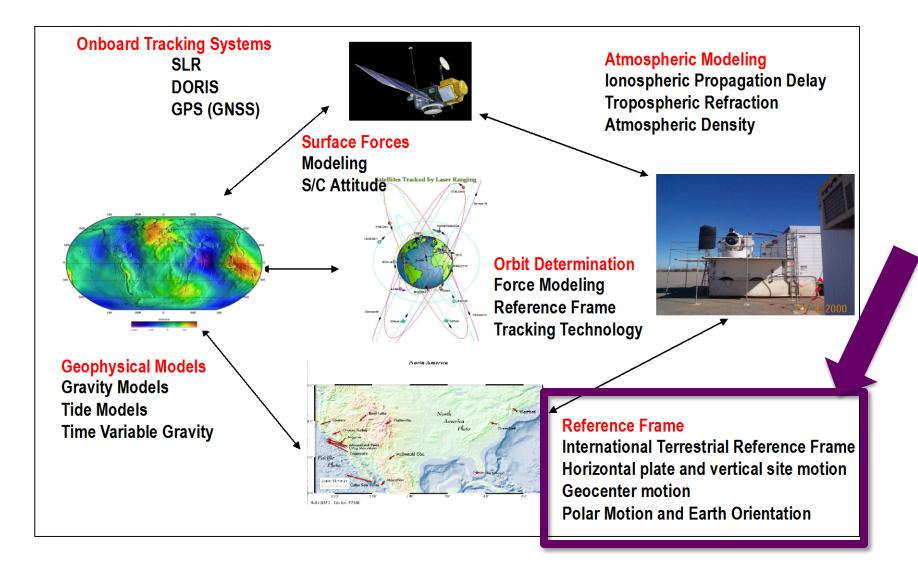
For TOPEX, we have this attitude information for 5-10% of the 10-day cycles over the entire mission – so we use an analytical model.



13:08:33 2015.04.04 Start global view X150



Orbit Determination Schematic





Reference Frame: What is it?

SLR



e.g. Hartebeesthoek (South Africa)



e.g. Greenbelt, Maryland

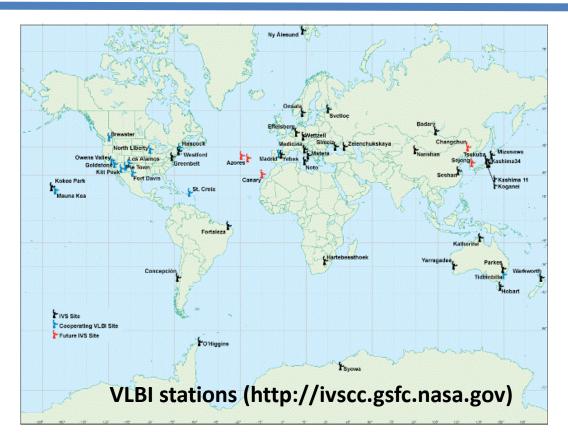
+ VLBI

╉





e.g. Greenbelt; Wettzell (Germany)



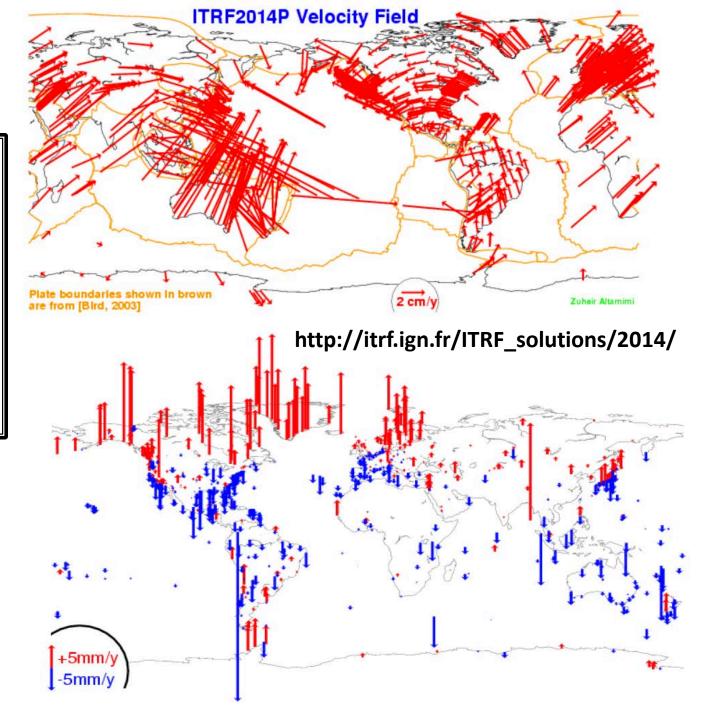




A reference frame realization consists of **positions and velocities of the reference points**.

For ITRF2014, postseismic relaxation is also modeled for the first time.

Figures from Zuheir Altamimi, IGN/France



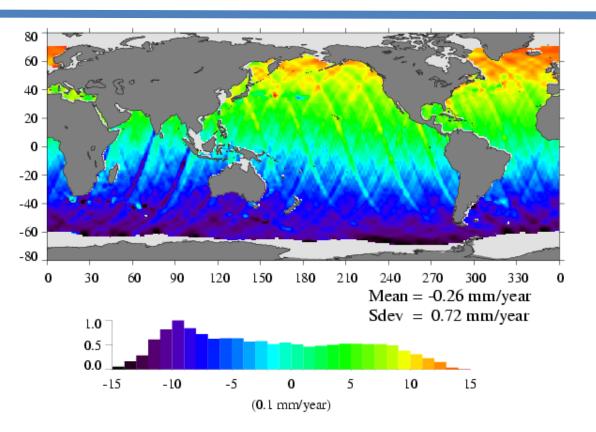


Altimeter satellites & TRF error

Reference Frame Realizations used in Altimeter Satellite POD.

- CSR95
- ITRF2000
- ITRF2005
- ITRF2008

Next. ITRF2014.

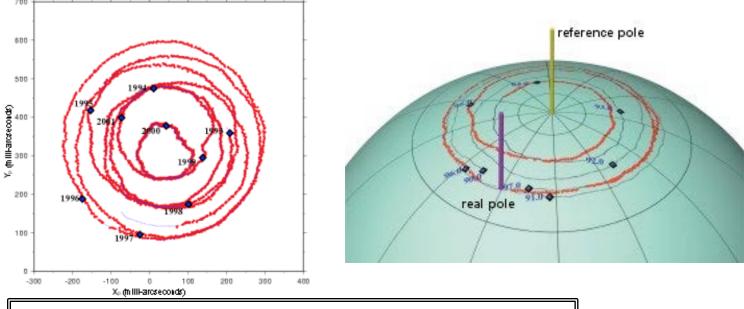


Regional **TOPEX** (1993-2002) Sea Surface Height Trend differences from direct impact of the ITRF2005 (GGM02C) minus CSR95 (JGM3) orbit differences. (from Beckley et al., *Geophys. Res. Lett.*, 2007).

Errors in the Z component of the TRF can produce large regional errors in MSL rate determination.



Polar Motion



1 arcsec = $1/3600^{\text{th}}$ of a degree; 0.1 arcsec = ~ 3 meters

The Earth rotates in an irregular fashion due to mass shifts and changes in its rotational velocity; The "true pole" can deviate by up to ten meters from the "reference pole". So on a regular basis, we must be supplied with updates which are provided by the IERS (International Earth Rotation and Reference Systems Service) – based on analysis of the space geodetic data (GPS, SLR, VLBI) that has been described.

Images Credit:.

Group de Recherche de Géodésie Spatiale (GRGS/CNES), Toulouse, France.

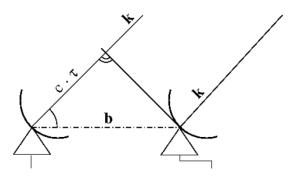


Altimetry Satellites, Earth Rotation, Quasars & VLBI



Active galactic nuclei, galaxies, quasars Distance 2 – 8 billion light years Point sources No proper motions

→quasi-inertial reference system



1 ms -> 7.4 cm on the Earth's equator

UT1 Determination Accuracy: 1972-1979: ~1 ms (Lunar Laser Ranging) 1979-1983: 0.4 ms (early VLBI) 1983-1991: 0.05 ms (campaign VLBI) (Feissel & Gambis, Adv. Space Res, 1993) VLBI is the only technique to measure the rotation "phase" of the Earth – how it changes with time.

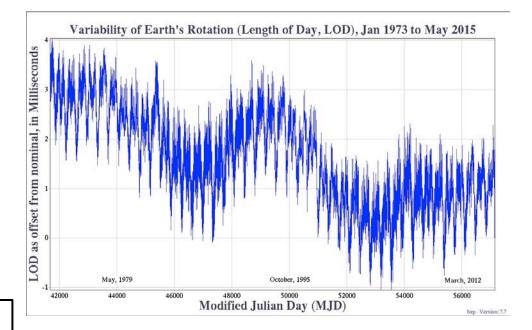


Image from US Naval Observatory, Earth Orientation Department

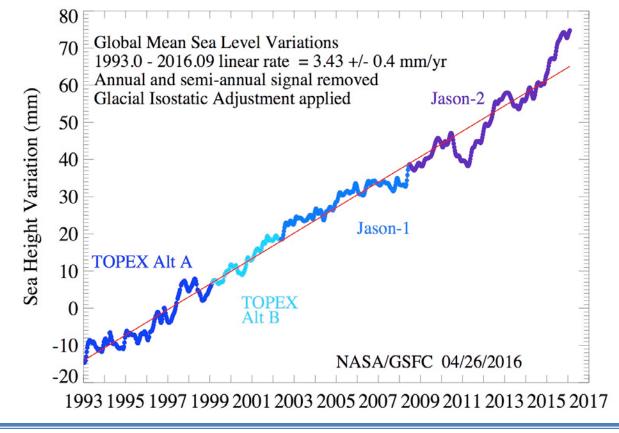


Application - Synoptic mapping of ocean height variations



Application – MSL determination

The precise orbits for TOPEX/Poseidon, Jason-1, Jason-2, all computed in a consistent reference frame (ITRF2008) are used to compute the global change in mean sea level from satellite ocean radar altimeter data.



http://podaac.jpl.nasa.gov/Integrated_Multi-Mission_Ocean_AltimeterData



Summary

• The Earth is a dynamic planet. In order to operate an observing system of altimeter satellites to monitor the global variations in ocean topography – we need to rely on precision tracking systems, detailed models of the forces that perturb the spacecraft orbit, and we must model in detail the observations, including target and propagation (media) effects.

- Inputs for Precise Orbit Determination Include
- (1) Terrestrial reference frame (updated every five years). ---
- (2) Precise model of Earth's gravity field including time variations determined from GRACE+GOCE data + supplemented by analysis of other satellite data.
- (3) Model the mass motions of the atmosphere \rightarrow So we rely on atmosphere models (ECMWF, NCEP) to account for these motions.
- (4) Polar motion and Earth rotation information (updated delay for near-real time products; every few weeks for higher precision products)